

Appendix A: Turbulence properties

Sub-grid turbulent diffusion (K_s) and the turbulent kinetic energy dissipation rate (ε)

The Smagorinsky's turbulence closure term [1], for the sub-grid turbulent diffusion (K_s) is calculated according to [2]:

$$K_s = (f_D C_s \Delta)^2 \sqrt{\frac{1}{2} \left(-\frac{\partial u_z}{\partial x} + \frac{\partial u_x}{\partial z} \right)^2 + \left(\frac{\partial u_x}{\partial x} \right)^2 + \left(\frac{\partial u_z}{\partial z} \right)^2} \quad (\text{Eq. A1})$$

Here, C_s is the Smagorinsky constant set to 0.2, Δ the size of the spatial discretization and u_z and u_x are the velocities in the vertical and horizontal dimension, respectively. The near sea surface turbulence shares many characteristics of that of a wall [3,4]. f_D describes the wall effect of the sea-surface on advection and diffusion ('law of the wall'), which is given by the van Driest wall damping function [5]:

$$f_D = 1 - \exp\left(-\frac{y^+}{A}\right)^2 \quad (\text{Eq. A2})$$

where A is the van Driest constant ($A = 26$) and y^+ the dimensionless distance from the sea surface according to $y^+ = \frac{u_\tau z}{\nu}$. Here ν is the kinematic viscosity ($1.05 \times 10^{-6} \text{ m}^2 \text{ s}^{-1}$ for seawater of 35 PSU and

20°C) and $u_\tau = \sqrt{\tau_w / \rho}$ is the frictional velocity scale, with ρ being the density of sea water and τ_w

the wall shear stress $\tau_w = \mu \sqrt{\left(\frac{\partial u_x}{\partial z}\right)^2}$, where μ is the dynamics viscosity set to $0.001 \text{ kg m}^{-1} \text{ s}^{-1}$ (for

seawater of 35 PSU and 20°C).

In free convective system and within the mixed layer the turbulent kinetic energy dissipation rate ε (W kg^{-1}) is equal to is the surface buoyancy flux J_0^b and remains relatively similar over the entire depth [4]. In such systems the vertical velocity scales according to $u_z = (J_0^b h_z)^{1/3}$ [6,7]. By replacing J_0^b with ε and solving for ε we can describe energy dissipation according to [4,8]:

$$\varepsilon = \frac{u_z^3}{h_z} \quad (\text{Eq. A3})$$

Using the mixing length time scale τ , with $\tau = 12$ hours [7], modelled values of ε ranging from $\sim 1.25 \cdot 10^{-10} \text{ W kg}^{-1}$ at $h_z = 100\text{m}$ to $\sim 5 \cdot 10^{-8} \text{ W kg}^{-1}$ $h_z = 2000\text{m}$ (figure A1).

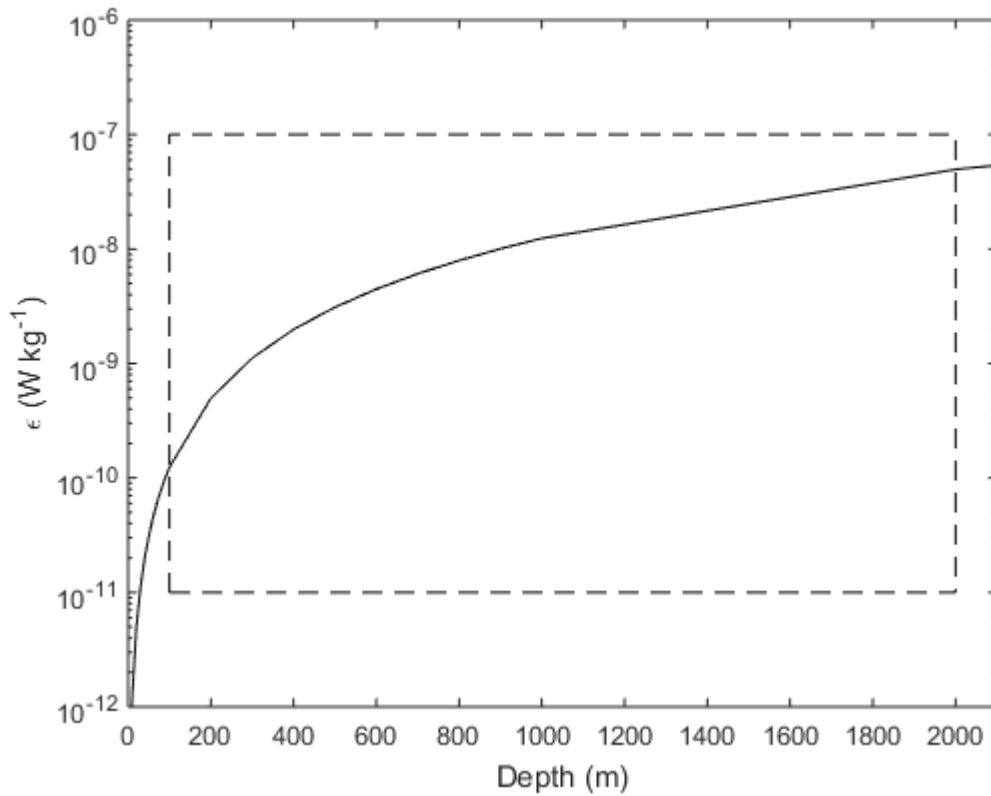


Figure A1: Modeled turbulent energy dissipation rate (ε) using the convective mixing length scale $\tau = 12$ hours [7]. The dashed box represents the approximate range of observed turbulent energy dissipation rates (ε) and vortex depth in deep convection [4,9–11].

Vertical profiles of modelled turbulent diffusivity (K_T)

The vertical profiles of the mean turbulent diffusivity (K_T) within the vortex show lower values at the sea surface, the highest values at around 30% of the depth and the slowest at the bottom (figure A2). These profiles are generally in agreement with those from application of LES used to simulate deep convection [12].

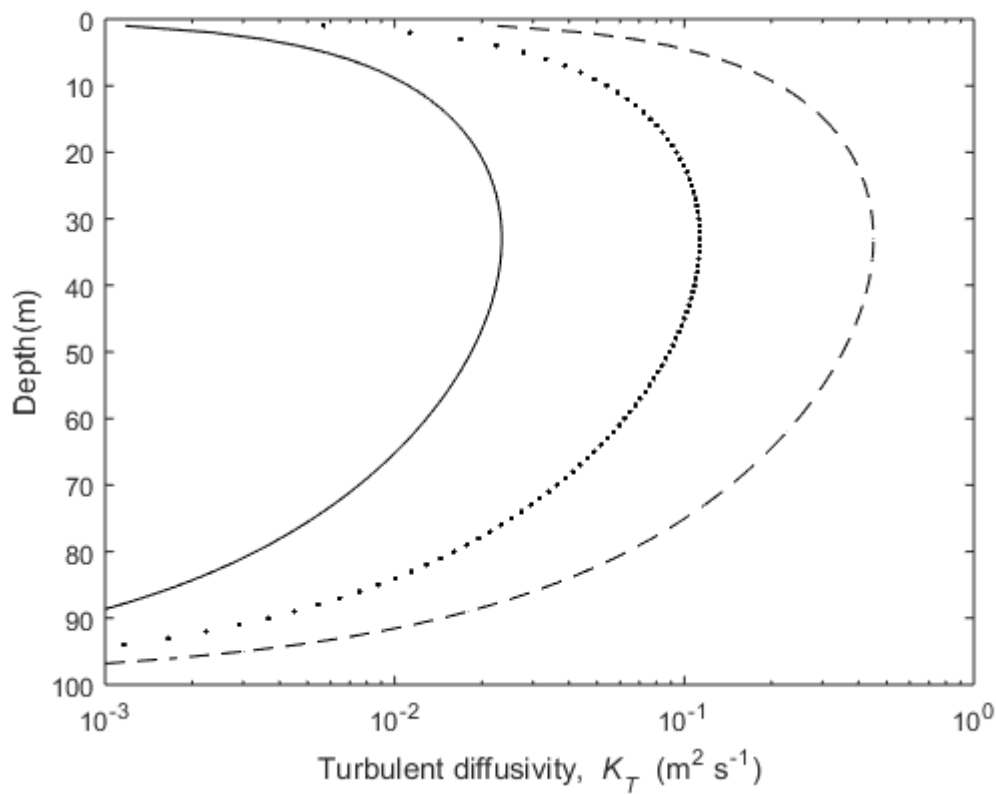


Figure A2: Profiles of turbulent diffusivity (K_T) for a 100 meter deep vortex with mean velocities of 25 $m d^{-1}$ (solid line), 75 $m d^{-1}$ (dotted line) and 200 $m d^{-1}$ (dashed line).

Spatial phytoplankton distribution with horizontally homogenous turbulent diffusivity (K_T)

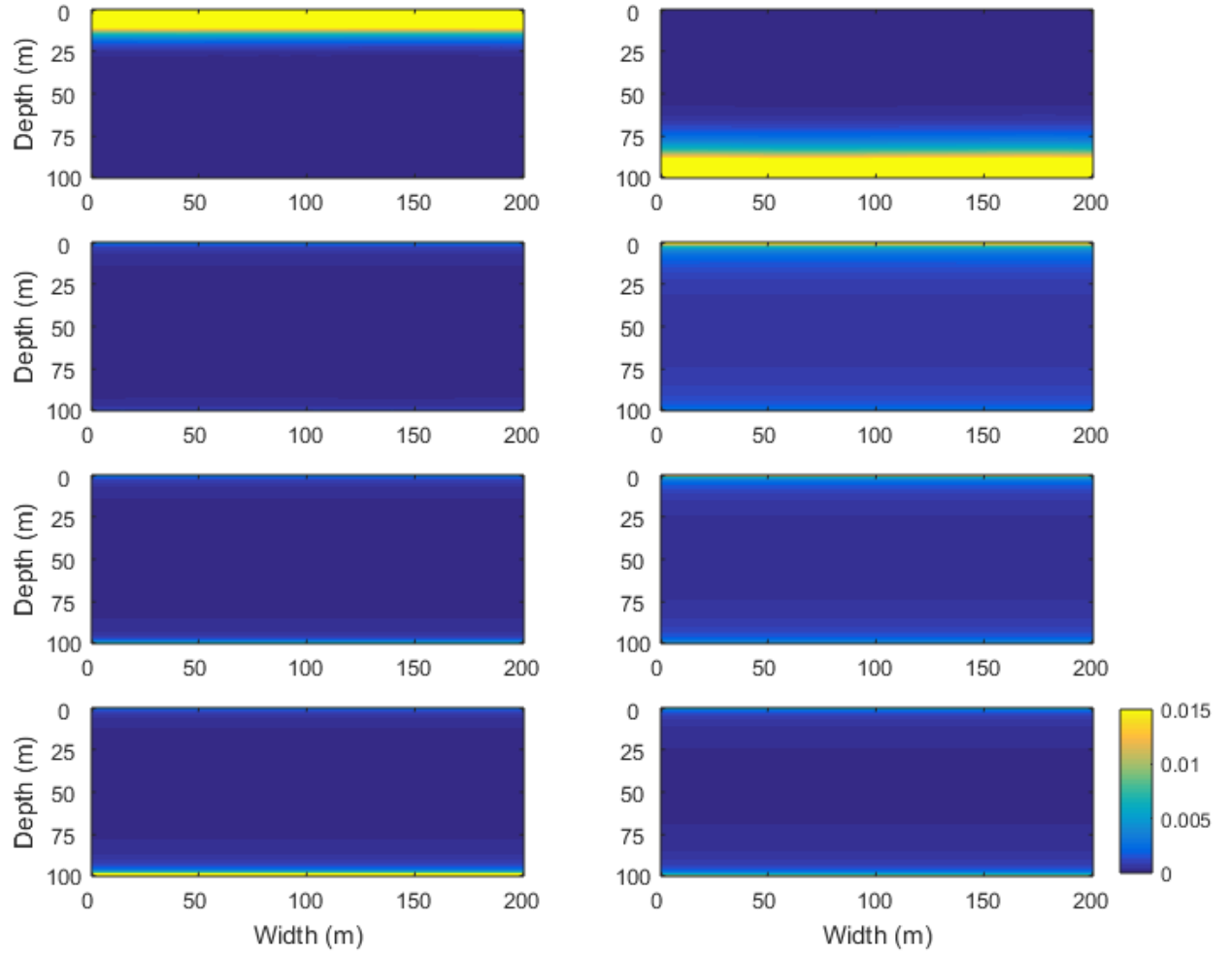


Figure A3: Spatial distribution (right eigenvector of the dominant eigenvalue, v_r) of neutrally buoyant (left column) and fast sinking (10 m d⁻¹, right column) phytoplankton. The maximum flow field velocities are (a,b) 0 m d⁻¹, (c,d) 50 m d⁻¹, (e,f) 200 m d⁻¹ and (g,h) 1000 m d⁻¹. The turbulent diffusivity (K_T) is parameterized according to Eq. 7 and the mean vertical profile of K_T is applied homogenously in the horizontal dimension.

Estimates of observed turbulent diffusivity (K_T)

In convective systems K_T can be described using the mean vertical velocity (\bar{u}) and the depth of the vortex (h_z) according to Eq. 5 [13].

In figure A4 the black dots show K_T estimated from observed values of \bar{u} and h_z , taken from [13–19]. Lavender et al. [16] found that peak vertical velocities are log normal distributed. Therefore, in convective systems where only the maximum vertical velocity (u_{\max}) was available the mean

vertical velocity was estimated according to $\bar{u} = 10^{\frac{\log_{10}(u_{\max})}{2}}$.

K_T can also be estimated from the surface buoyancy flux J_0^B , according to [13] $K_T = h_z^{3/4} |J_0^B|^{1/3}$.

In figure B4 the gray dots show K_T estimated from observed values of J_0^B . In convective system the surface buoyancy flux is related to the net surface heat flux J_0^h (W m^{-2}) through

$$J_0^B = \frac{\alpha g}{\rho_0 c_p} J_0^h, \text{ with } \alpha (2.5 \cdot 10^{-4} \text{ } ^\circ\text{C}^{-1}) \text{ being the thermal expansion coefficient; } g (9.81 \text{ m}^2 \text{ s}^{-1})$$

being the gravitational acceleration; ρ_0 (1037 kg m^{-3}) being the seawater density and c_p ($3998 \text{ J (kg}^\circ\text{C)}^{-1}$) being the heat capacity.

Hence, K_T can also be described as a function of J_0^h [20]:

$$K_T = h_z^{3/4} \left| \frac{\alpha g}{\rho_0 c_p} J_0^h \right|^{1/3} \quad (\text{Eq. A4})$$

In figure A4 the lines indicate theoretical estimates of K_T for different values of J_0^h .

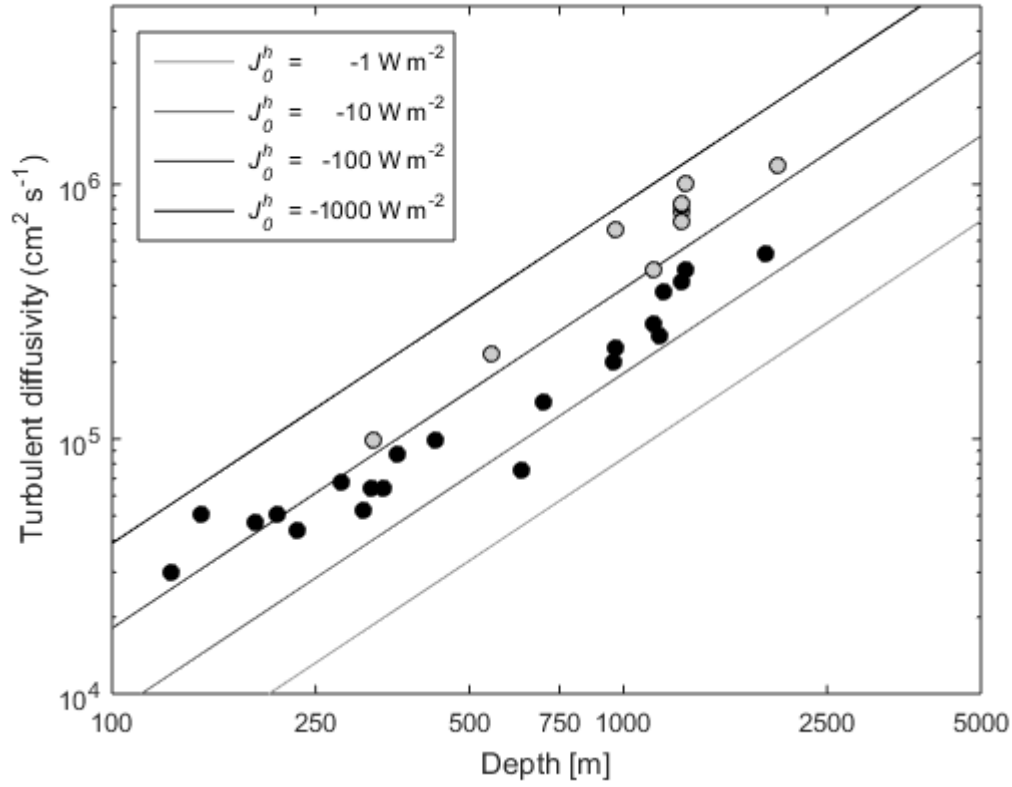


Figure A4: Observed and theoretical estimates of turbulent diffusivity K_T . The black dots are estimated from vertical velocities (Eq. 5), the gray dots are estimated from surface buoyancy flux J_0^B and the lines are theoretical estimates (using Eq. A4) for J_0^h with values -1, -10, -100 and -1000 W m^{-2} .

Net population growth rate and light extinction coefficients

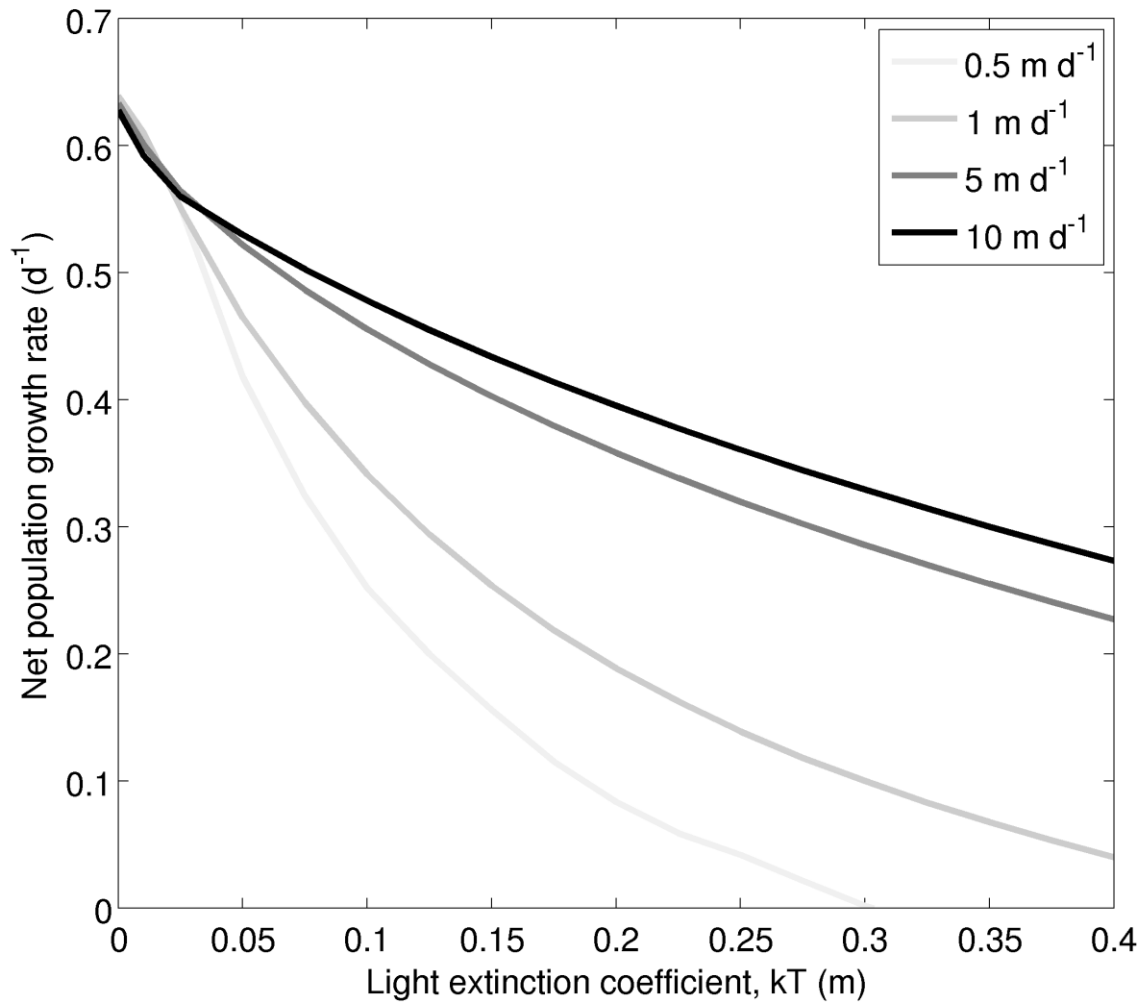


Figure A5: Changes of the net population growth rate with different light extinction coefficients at different levels of turbulent advective transport. The depth of the domain is 100 m. All other parameters are given in table 1.

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